

Written and collected by Zia H Shah MD

# Mountain Range Formation: Timescales and Geological History

## Mountain-Building Timeframes – Forming a Mountain Range

Mountain ranges do not form overnight – they typically develop over **tens of millions of years** or longer as tectonic forces slowly crumple and uplift the crust.

A mountain-building event (an **orogeny**), usually spans a vast timescale: **orogenies often take place over tens or even hundreds of millions of years**<sup>1</sup>. For example, the **Himalayan** range began rising around **50 million years ago** when India collided with Asia and **continues to rise today**<sup>2</sup>. In the case of the **Andes**, subduction of the Nazca plate under South America initiated uplift tens of millions of years ago; recent studies suggest the Andes reached their towering heights in stages, with a major pulse of rapid uplift occurring only **~10–6 million years ago** after long-term crustal thickening<sup>3</sup>. Similarly, the **Rocky Mountains** of North America formed during the Laramide Orogeny, which **lasted roughly 40 million years** (approximately **75–35 Ma**)<sup>4</sup>. These examples illustrate that a “single” mountain range can take on the order of **10–100+ million years** to fully develop from the onset of tectonic collision to peak elevation. The prolonged timeframe arises because convergence and uplift happen incrementally: plates collide at only a few centimeters per year, and mountains continue rising until tectonic forces wane or erosion counterbalances uplift.

## Earth’s First Mountains – Early Orogeny in Deep Time

Our planet is about **4.54 billion years old**<sup>5</sup>, and once the early crust solidified and plate tectonics began, mountain-building processes would have started. Geological evidence indicates that **mountain ranges were already forming by the Archean Eon** (over 3 billion years ago). One of the **oldest known mountain remnants** is the Barberton Makhonjwa **Greenstone Belt** in South Africa, which dates to approximately **3.5 billion years ago**<sup>6</sup>. This shows that by 3.5 Ga (billion years ago), volcanic activity and continental collisions were uplifting Earth’s crust into proto-mountain ranges. By the **late Archean (~2.7–2.5 Ga)**, major orogenic events like the **Kenoran (Algoman) orogeny** were occurring; this event – a series of continent collisions in what is now North America – lasted on the order of <100 million years and **brought the Archean Eon to a close around 2.5 Ga**<sup>7</sup>. The end of the Archean marks a significant shift in Earth’s crustal development: by ~2.5 Ga, the first stable continental cores (cratons) had formed, and continental crust had thickened to ~40 km, indicating substantial mountain-building had taken place<sup>8</sup>. In summary, **mountain ranges began forming as early as ~3.5 billion years ago**, not long after Earth’s crust became stable, demonstrating that plate tectonic and orogenic processes have been active for most of Earth’s history.

## Major Orogenic Events: A Geological Timeline

Earth’s geologic record preserves numerous **orogenies (mountain-building events)**. Below is a timeline of **major mountain range formation events** throughout Earth’s history, with approximate ages and notable examples:

- **~3.5 billion years ago (Archean Eon):** *Earliest mountain building. Greenstone belts* like the Barberton Mountains (South Africa) formed ~3.5 Ga<sup>6</sup>, indicating the early presence of mountains on the young Earth. These ancient

ranges were the result of **volcanic island arcs and microcontinents accreting** onto early cratons, providing the first evidence of orogeny in Earth's history.

- **~2.7–2.5 billion years ago (Late Archean): Kenoran (Algoman) Orogeny.** Multiple continental blocks (e.g. the Superior Province in Canada) collided and sutured together during this period <sup>7</sup>. The Kenoran orogeny lasted on the order of tens of millions of years and culminated ~2.5 Ga, marking the assembly of major parts of the Canadian Shield and the completion of Archean continental crust formation. By this time, Earth's continents had grown significantly through mountain-building – the Archean-Proterozoic transition saw the crust reach near-modern thickness <sup>8</sup>.
- **~2.2–2.0 billion years ago (Paleoproterozoic): Early Proterozoic orogenies.** Orogenic events continued as the first supercontinents took shape. For example, the **Eburnean orogeny** in West Africa (c. 2.2–2.0 Ga) was a major mountain-building episode during the formation of the ancient supercontinent **Nuna/Columbia** <sup>9</sup>. Around ~1.9–1.8 Ga, the **Trans-Hudson Orogeny** (North America) similarly sutured Archean cratons together, contributing to Nuna's assembly. These events created extensive mountain belts across what would become the cores of modern continents.
- **~1.3–1.0 billion years ago (Mesoproterozoic): Grenville Orogeny – Rodinia Supercontinent.** A long-lived series of collisions occurred as continents amalgamated into the supercontinent **Rodinia**. The **Grenville orogeny** (peaking ~1.1 Ga) was a worldwide mountain-building event associated with Rodinia's assembly <sup>10</sup>. For instance, eastern North America underwent massive crustal compression when the continent **Laurentia collided with Amazonia (~1.1 Ga)**, raising a huge mountain range – the eroded roots of which are exposed today in the Appalachians and Canada <sup>11</sup>. Grenville-age belts are found on several continents, attesting to a global orogenic episode. This orogeny lasted many tens of millions of years (c. 1250–980 Ma) <sup>12</sup> and created some of the largest mountains of the Proterozoic Eon.
- **~650–500 million years ago (Neoproterozoic to Early Paleozoic): Pan-African Orogeny and Gondwana Assembly.** As Rodinia broke apart, new collisions gave rise to Gondwana. The **Pan-African orogeny** (a series of mountain-building events ~650–530 Ma) stitched together Africa, South America, Antarctica, India, and Australia into the supercontinent **Gondwana** <sup>13</sup>. One phase, the **East African Orogeny (~750–550 Ma)**, involved the collision of East and West Gondwanan fragments. These events built huge ranges in what are now Africa and Antarctica. By the dawn of the Cambrian (~541 Ma), Gondwana's formation was complete, accompanied by extensive highlands (though many have since eroded away).
- **~480–400 million years ago (Ordovician–Devonian Periods): Taconic, Caledonian, and Acadian Orogenies.** During the Paleozoic, numerous collisions occurred along the margins of the ancient continent Laurentia (North America) and Baltica. Around **470 Ma**, the **Taconic orogeny** occurred as an island arc collided with eastern North America, initiating the formation of the **Appalachian Mountains** <sup>14</sup>. Later, as the **Iapetus Ocean** continued to close, the **Caledonian orogeny** (~430–400 Ma) joined parts of Europe (Baltica) with Laurentia (forming the Caledonide Mountains in Scotland, Scandinavia, and Greenland – the ancient continuation of the Appalachians). By **~380 Ma**, the **Acadian orogeny** further uplifted the Appalachians (particularly in the northern Appalachians) <sup>15</sup>. These combined events raised a vast mountain chain spanning what is now North America and Europe.
- **~360–300 million years ago (Carboniferous Period): Variscan (Hercynian) Orogeny.** In the late Paleozoic, as multiple landmasses converged, the **Variscan orogeny** (also called Hercynian) built mountains across what is now Europe (e.g. the Massif Central in France, Armorican and Bohemian highlands). This occurred ~340–300 Ma when Gondwana (Africa-South America) collided with Laurussia (Europe-North America) from the south. The Variscan belts in Europe were contemporaneous with the final assembly of Pangaea and correspond to the **Alleghanian orogeny** in North America.

- **~320–250 million years ago (Pennsylvanian–Permian Periods): Alleghanian Orogeny – Formation of Pangaea.** The culmination of the Paleozoic orogenies came when **all major continents merged into the supercontinent Pangaea**. In North America, the **Alleghanian orogeny (~320–300 Ma)** occurred as ancestral Africa (part of Gondwana) collided with Laurentia, **uplifting the central and southern Appalachians** to Himalayan-scale heights <sup>16 17</sup>. By ~290 Ma, the collision of Africa and North America had closed the **Iapetus Ocean** <sup>18</sup>. Around the same time, Eurasia was completed by the collision of Siberia with Europe (the **Uralian orogeny**, Permian period), forming the **Ural Mountains**. By ~250 Ma (end of the Permian), **Pangaea** was fully assembled, and extensive mountain ranges (Appalachians, Urals, Variscides, etc.) stood tall along the sutures of the supercontinent.
- **~200–100 million years ago (Mesozoic Era): Breakup of Pangaea and Mesozoic Orogenies.** The Mesozoic saw Pangaea split apart, but mountain-building continued along active plate margins. For example, as North America overrode the Pacific plates, **Cordilleran orogenies** built the western mountain systems. The **Sonoma orogeny** (Permian/Triassic, ~250 Ma) and later the **Sevier orogeny** (Jurassic–Cretaceous, peaking ~150–100 Ma) created early versions of the **Rocky Mountains** and ranges farther west. These involved accretion of island arcs and crustal shortening in what is now the U.S. Cordillera. Toward the end of the Mesozoic, the **Laramide orogeny** (Late Cretaceous to Paleogene) uplifted the Rocky Mountains inland. The Laramide orogeny **lasted ~40 Myr (75–35 Ma)** <sup>19</sup> raising the high peaks of today's Rockies. By the time the dinosaurs went extinct (~66 Ma), the Rockies were well underway, though mountain-building in the American West persisted into the Eocene (~35 Ma) as the Laramide waned.
- **~65 million years ago to present (Cenozoic Era): Alpine–Himalayan Orogeny (Still Ongoing).** In the Cenozoic, continental collisions associated with the African and Indian plates created a vast interconnected mountain belt from Europe through Asia. The **Alpine orogeny** (starting ~65–55 Ma) occurred as the African plate and smaller microplates pressed into Eurasia, raising the **Alps** and Carpathians (with major uplift ~30–5 Ma). Meanwhile, to the east, the **Indian Plate collided with Eurasia around 50 Ma**, initiating the **Himalayan orogeny** <sup>19</sup>. The **Himalayas** have been rising ever since – even today the range is actively growing due to ongoing convergence <sup>2</sup>. This Cenozoic mountain-building produced Earth's highest peaks (Mt. Everest and K2 in the Himalaya). It is part of a larger zone of collision – sometimes called the **Persia-Tibet-Burma orogeny** – that extends from the Mediterranean (where the Atlas and Zagros Mountains are uplifted) through the Middle East, Himalayas, and into Southeast Asia <sup>20</sup>. *Importantly, many of these Cenozoic orogenies are still active*, meaning mountain ranges like the Himalayas and Andes are geologically “young” and continue to evolve today.

**In summary**, mountain range formation is a slow but relentless process. Individual ranges usually require **tens of millions of years** of tectonic convergence and crustal deformation to rise to full prominence <sup>1</sup>. Earth's first mountains appeared over **3 billion years ago**, once the planet's surface had cooled enough for rigid plates to form <sup>6</sup>. Since then, Earth's history has been punctuated by numerous **orogenic events** – from the ancient cratonic collisions of the Precambrian to the grand collisions forming supercontinents like Pangaea and the ongoing Alpine-Himalayan system. These events, occurring at different times over the past **4.5 billion years**, have built and rebuilt the world's great mountain ranges. Mountains thereby serve as enduring records of plate tectonics through deep time, each range preserving the story of its uplift in the face of continual erosion and the ceaseless movements of Earth's crust.

**Sources:** The Geological Society of London <sup>2</sup>; USGS (This Dynamic Earth) <sup>19</sup>; *CK-12 Earth Science* (LibreTexts) <sup>1</sup>; Ultimate Kilimanjaro (Geological history of Barberton) <sup>6</sup>; *Wikipedia: Orogeny* <sup>4</sup>, Appalachian Mountains <sup>11</sup>, Algoman (Kenoran) orogeny <sup>7</sup>; University of Bristol News <sup>3</sup>; *TeacherFriendly Guide to Geology* (Paleontological Research Institution) <sup>14 16</sup>; and other geologic time scale references.

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